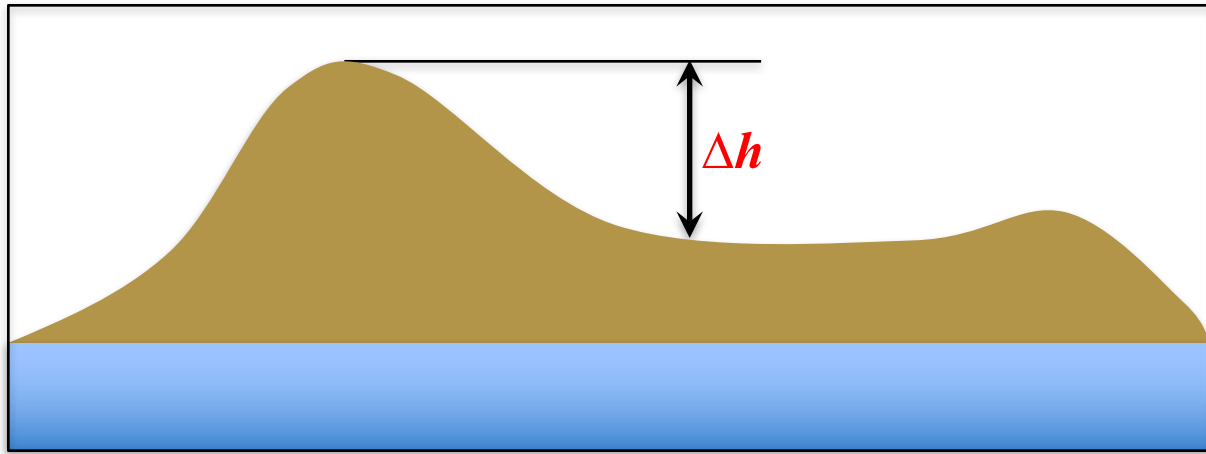


❖ It appears that under their own units the gravity gradient is a million times worse than the gravity.

❖ In reality this is not the case.

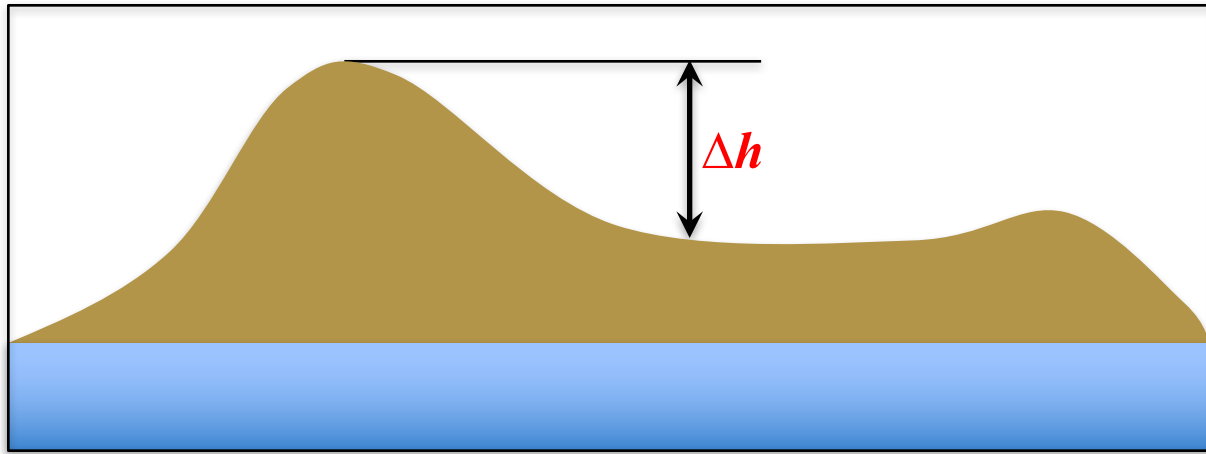
❖ In particular, the opposite is true in terms of dealing with the nuisance of elevation corrections.



$R_E = 6.371 \times 10^6$ meter, The average radius of the Earth

$$\Delta g = -1.9 \times 10^9 \mu gal \left(\frac{\Delta h}{R_E} \right)$$

$$\Delta \left(\frac{\partial g}{\partial z} \right) = -9.0 \times 10^3 E \left(\frac{\Delta h}{R_E} \right)$$



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The gradient has \sim Million times smaller elevation corrections than the gravity (under their own units)

Contents

- 1. Static mass anomaly with gravity gradient (a brief review)**
- 2. Deformation related gravity gradients (Preliminary modeling)**
 - Traditional gas deposit and production
 - Hydraulic cracking for shale gas

Gravity gradient tensor

$$\mathcal{G} = \begin{pmatrix} \frac{\partial g_1}{\partial x_1} & \frac{\partial g_1}{\partial x_2} & \frac{\partial g_1}{\partial x_3} \\ \frac{\partial g_2}{\partial x_1} & \frac{\partial g_2}{\partial x_2} & \frac{\partial g_2}{\partial x_3} \\ \frac{\partial g_3}{\partial x_1} & \frac{\partial g_3}{\partial x_2} & \frac{\partial g_3}{\partial x_3} \end{pmatrix}$$

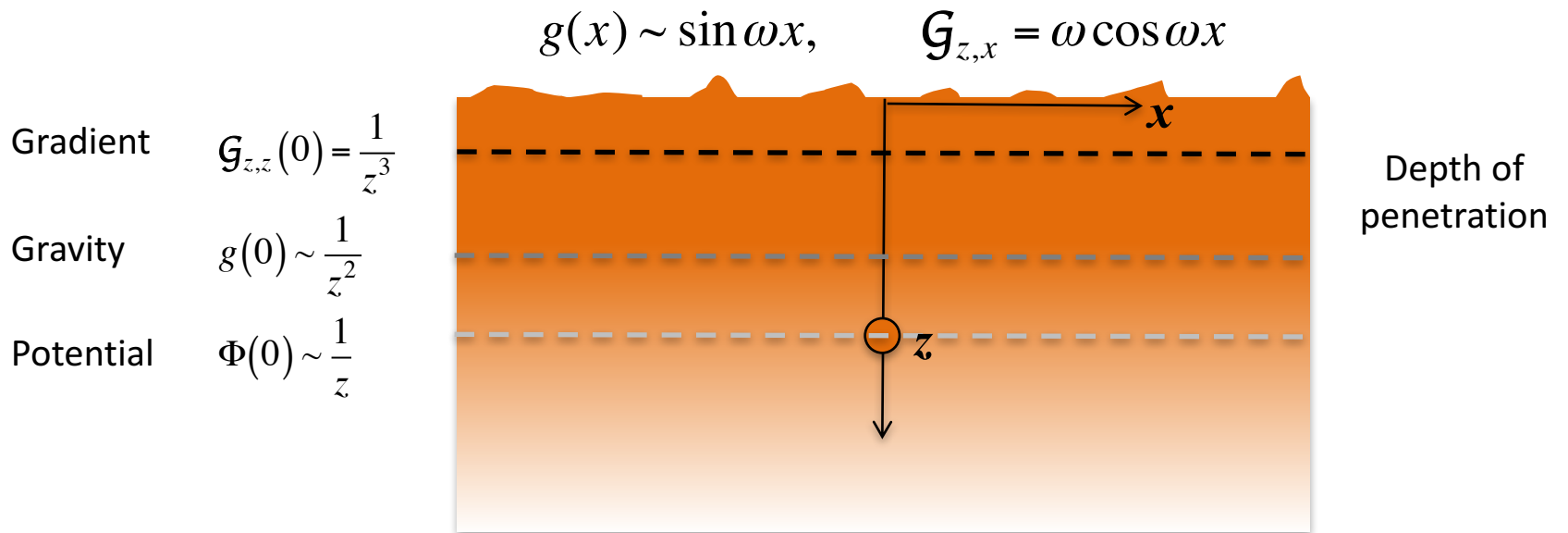
Symmetric

$$\mathcal{G}_{i,j} = \frac{\partial^2 \Phi}{\partial x_i \partial x_j}$$

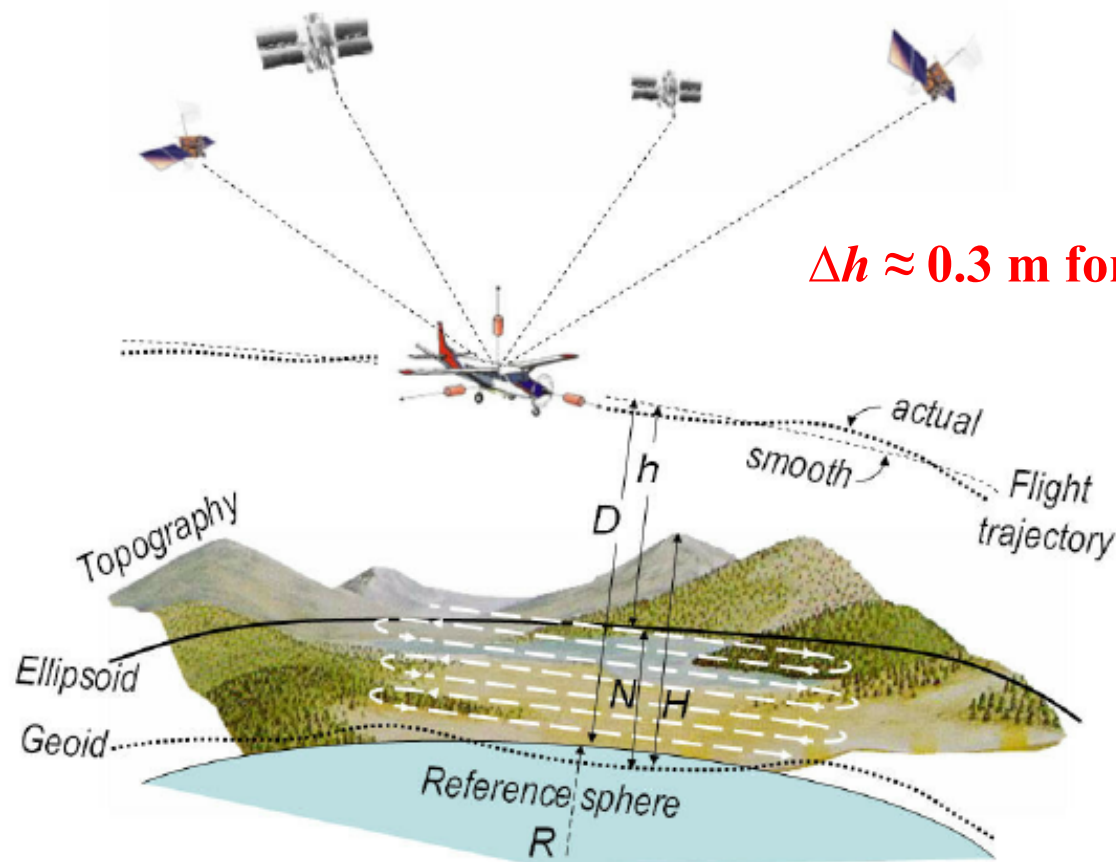
Zero trace

$$\sum_{i=1}^3 \mathcal{G}_{i,i} = \sum_{i=1}^3 \frac{\partial^2 \Phi}{\partial x_i^2} = 0$$

Gradient is sensitive to shallow and sharp structures

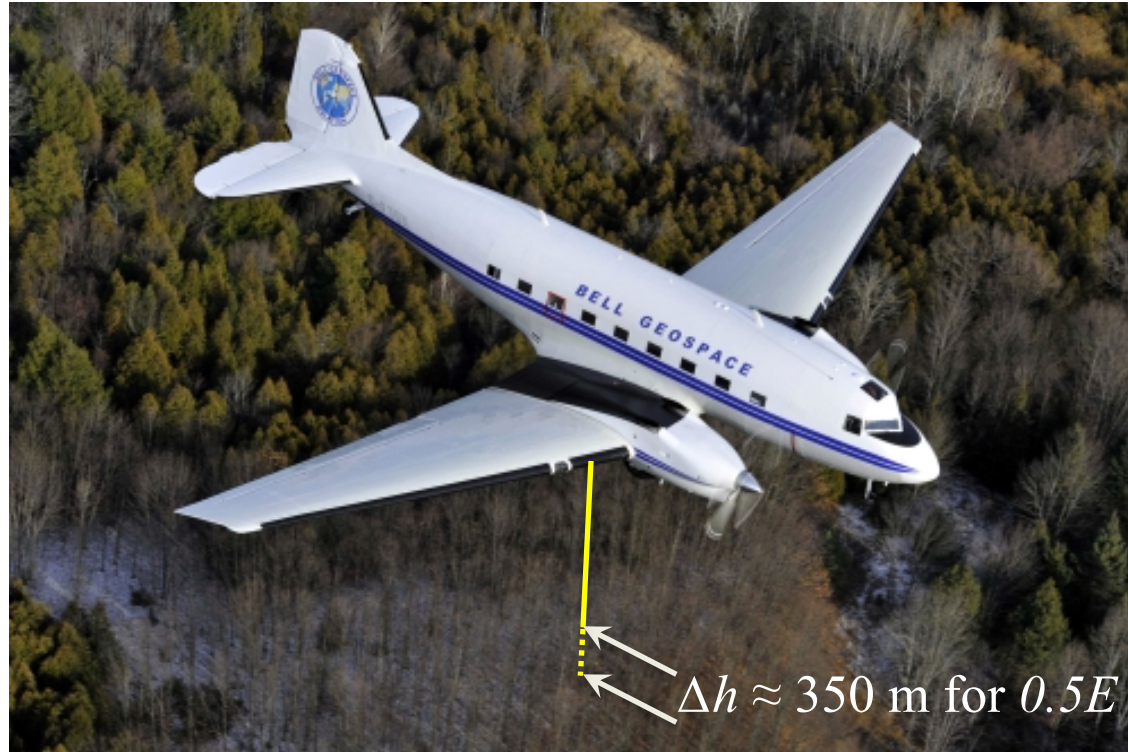


A big nuisance for airborne gravity is the height correction for repeated flyby



$\Delta h \approx 0.3 \text{ m for } 0.1 \text{ mgal}$

For airborne gravity gradient the accuracy of aircraft height is much less important



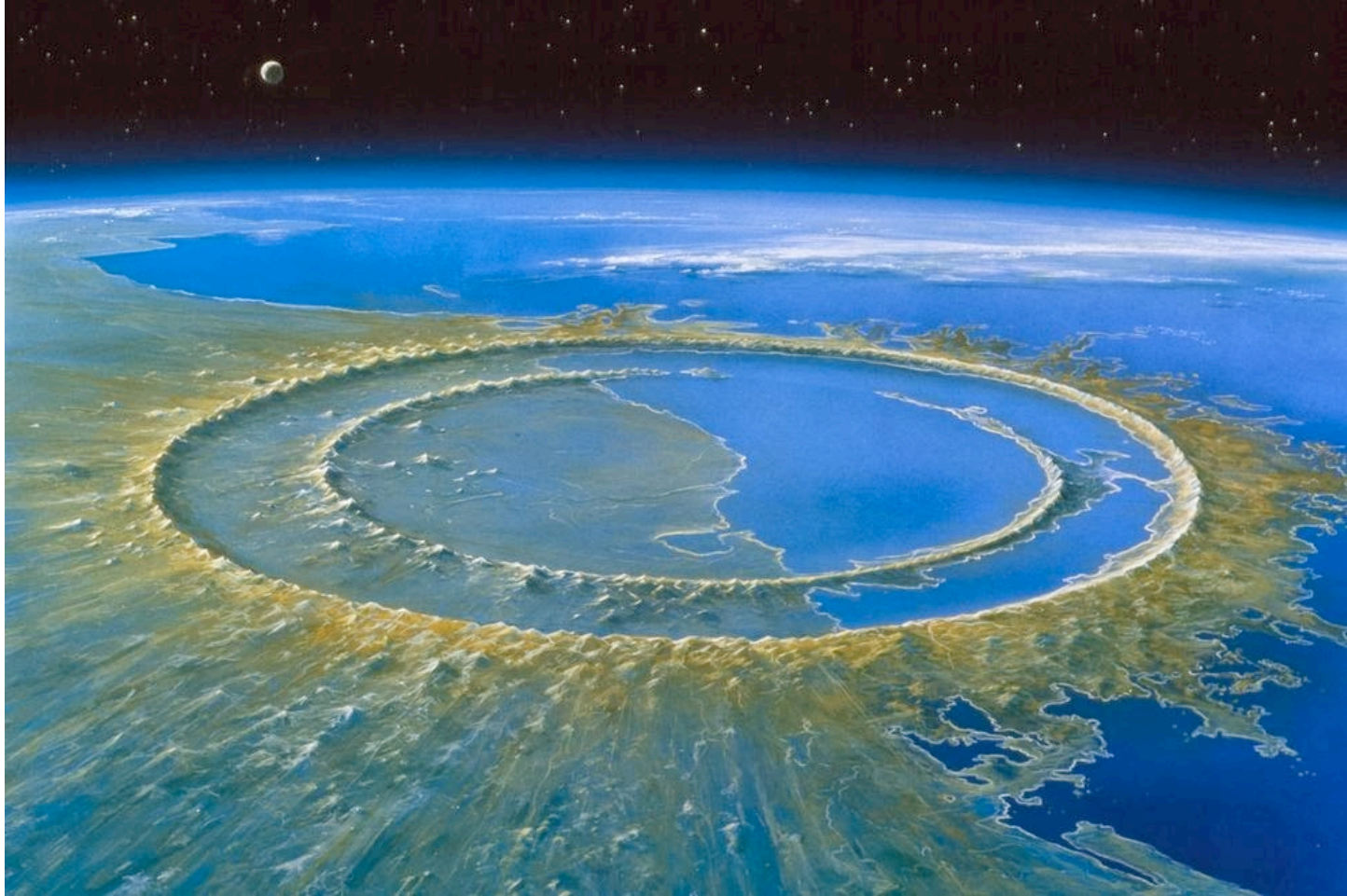
<http://www.intrepid-geophysics.com/ig/index.php?page=gradiometry>

Probe the static structure: the double rings of Vredefort Crater, South Africa, with Gravity Gradient



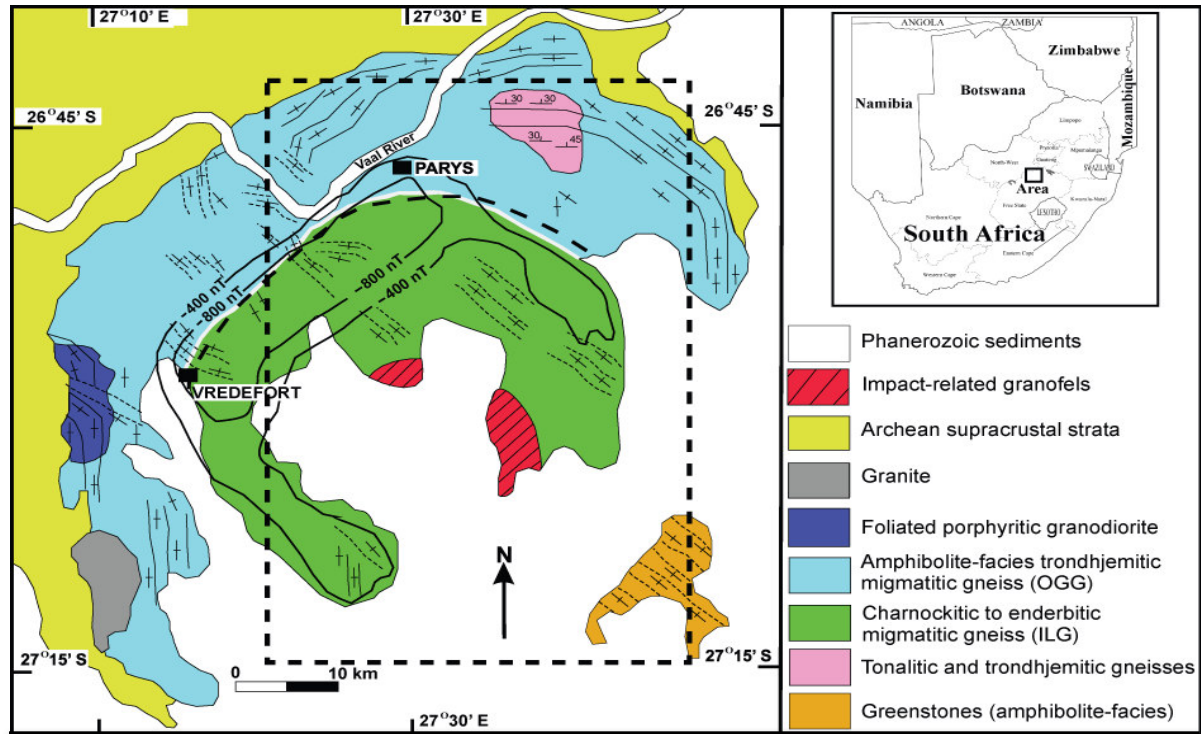
https://en.wikipedia.org/wiki/Vredefort_crater#/media/File:Vredefort_crater.jpg

Artistic Reconstruction of the Crater from geological evidence



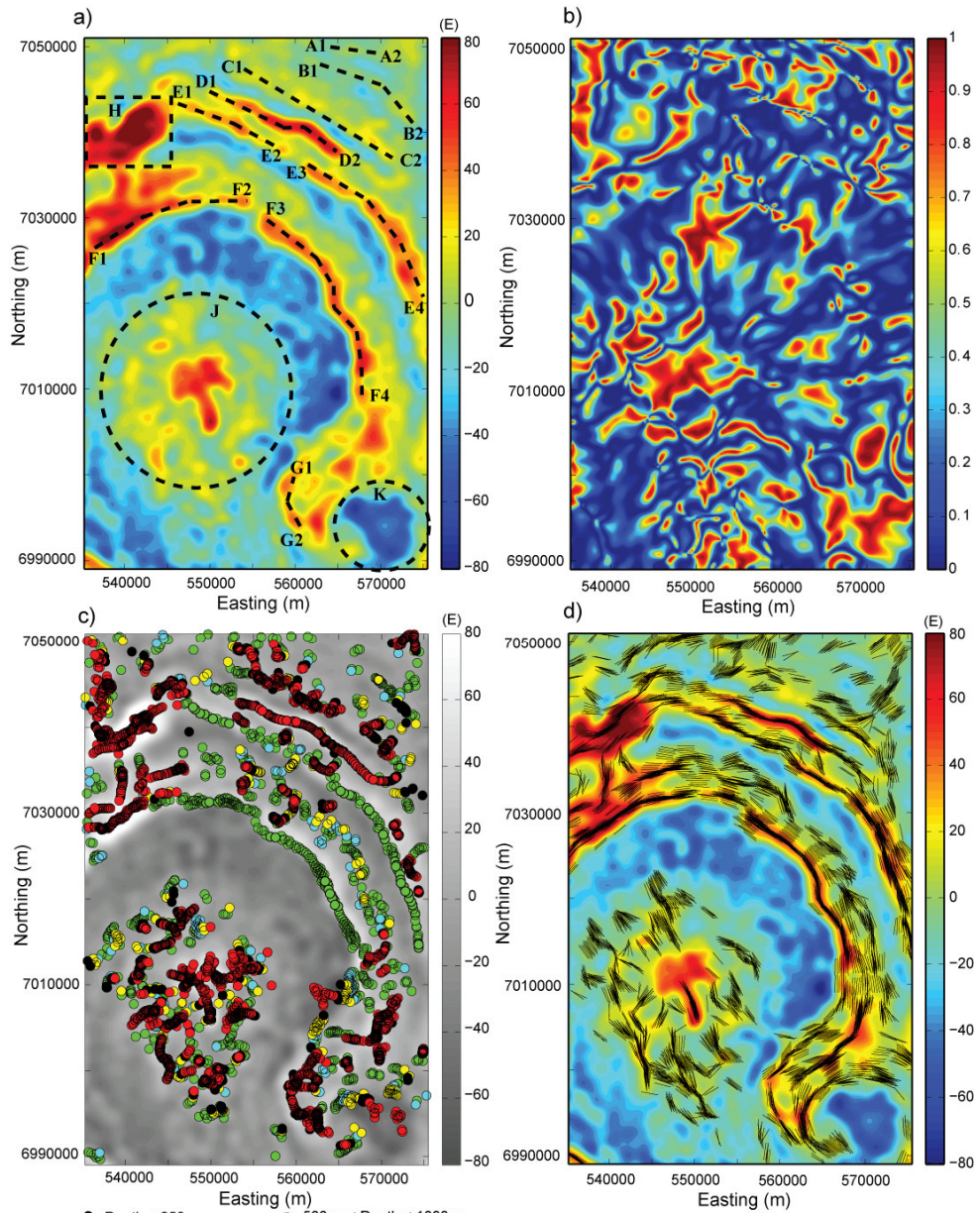
<http://news.nationalgeographic.com/news/2013/13/130214-biggest-asteroid-impacts-meteorites-space-2012da14/>

Investigating area by airborne gradiometer



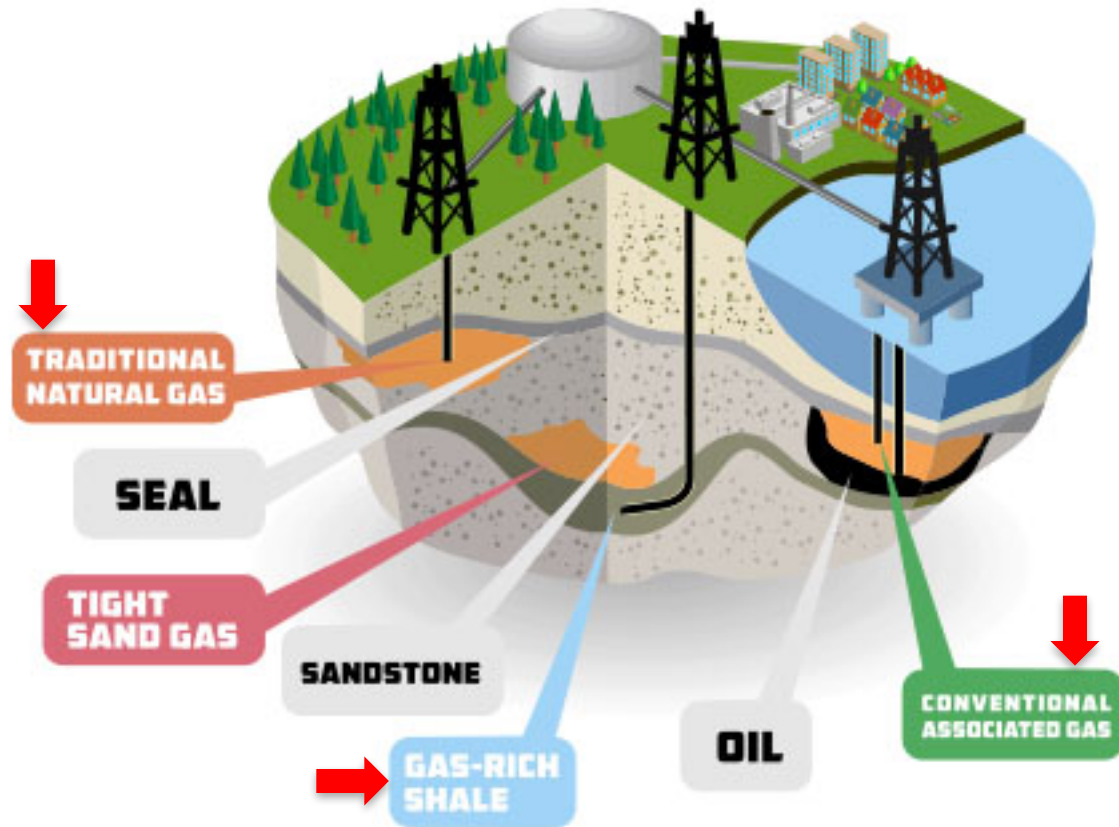
Beiki *et al.* (2010)

$$\frac{\partial g_z}{\partial z}$$



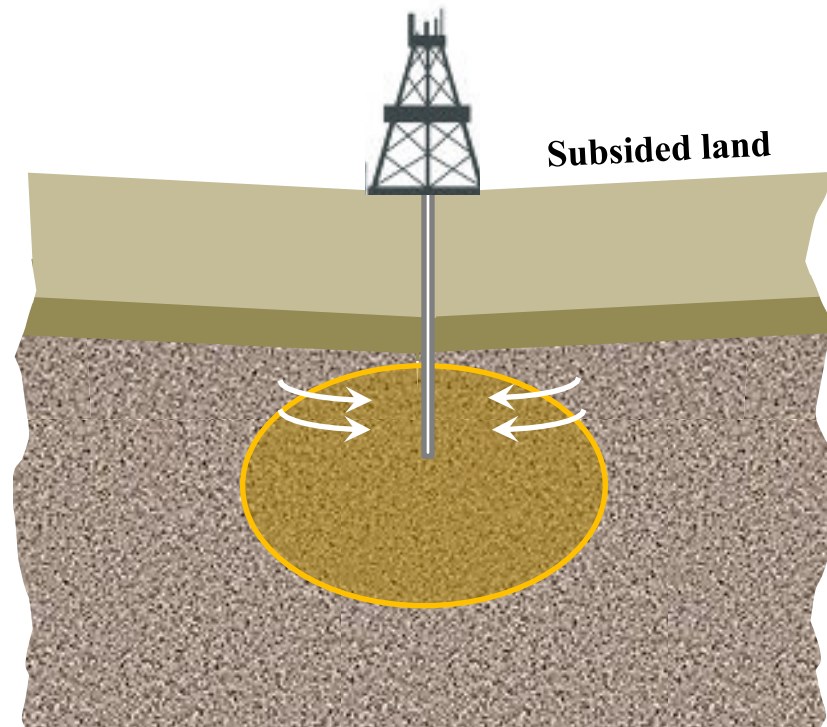
Beiki *at al.* (2010)

Probe natural gas production



<https://www.waltongas.com/index.php/blog/category/44/types-of-natural-gas/>

Traditional natural gas production

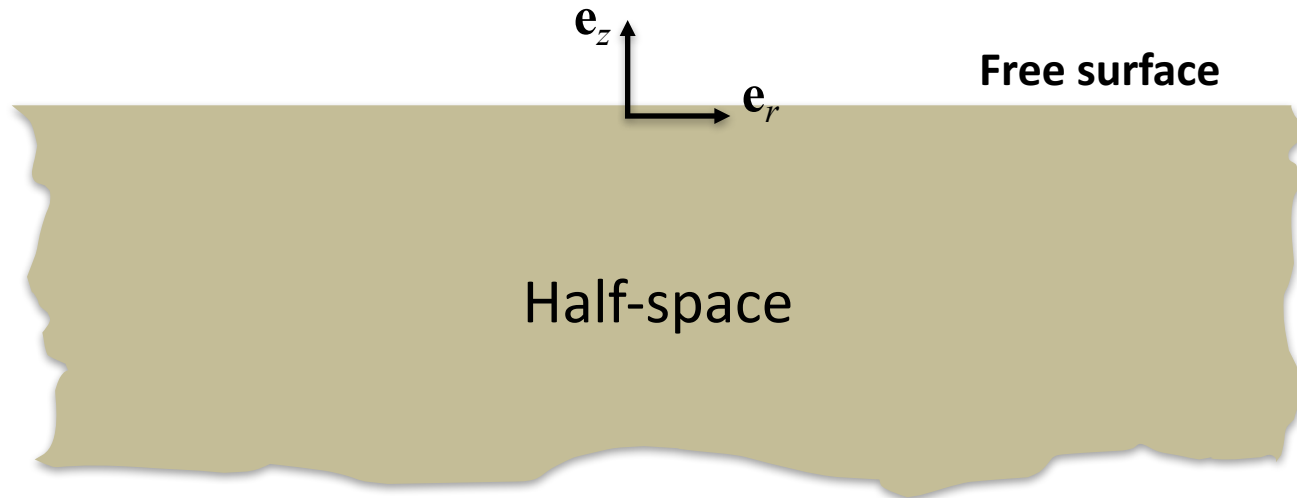


Gavity can be used to monitor the underground mass changes. We will show that the gravity signal is mostly dominated by the land subsidence /uplift, while the gravity gradient is land-subsidence/uplift proof.

Self-gravitating elastic half space model set up

\mathbf{u}_h ----- Displacement

ϕ ----- Perturbed gravitational potential



$$\nabla^2 \mathbf{u}_h + \frac{1}{1-2\sigma} \nabla \nabla \cdot \mathbf{u}_h - \frac{\rho_0}{\mu} \nabla \phi + \frac{\rho_0 g}{\mu} (\mathbf{e}_z \cdot \nabla \mathbf{u}_h - \mathbf{e}_z \nabla \cdot \mathbf{u}_h) = 0$$

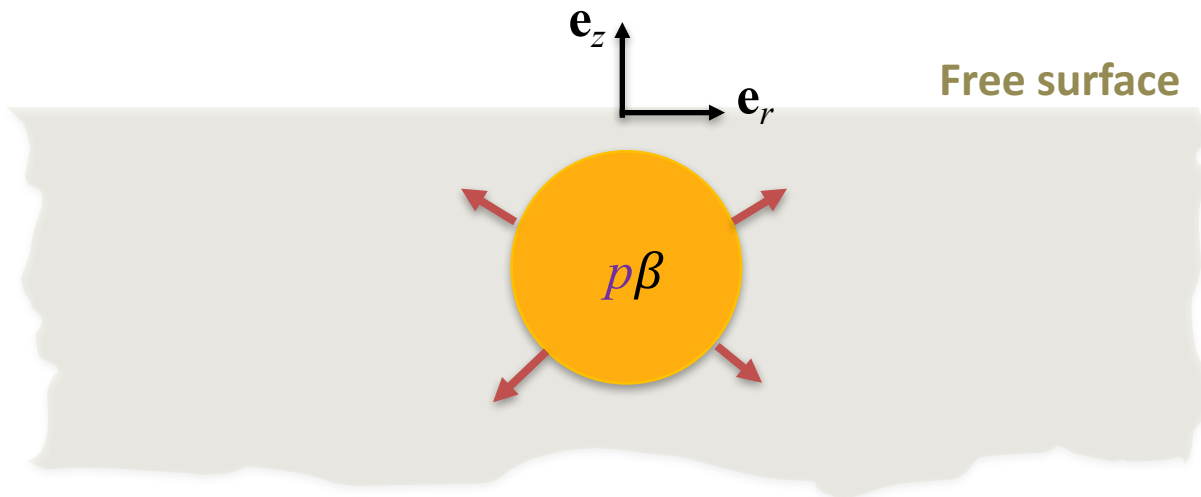
$$\nabla^2 \phi = -4\pi G \rho_0 \nabla \cdot \mathbf{u}_h$$

$$\lim_{\substack{R \rightarrow \infty \\ z \rightarrow +\infty}} (\mathbf{u}_h, \phi) \rightarrow 0$$

Poroelastic ball embedded in a self-gravitating half space

p ----- Incremental pressure variation

β ----- Poroelastic expansion coefficient



Decoupled pore mechanics

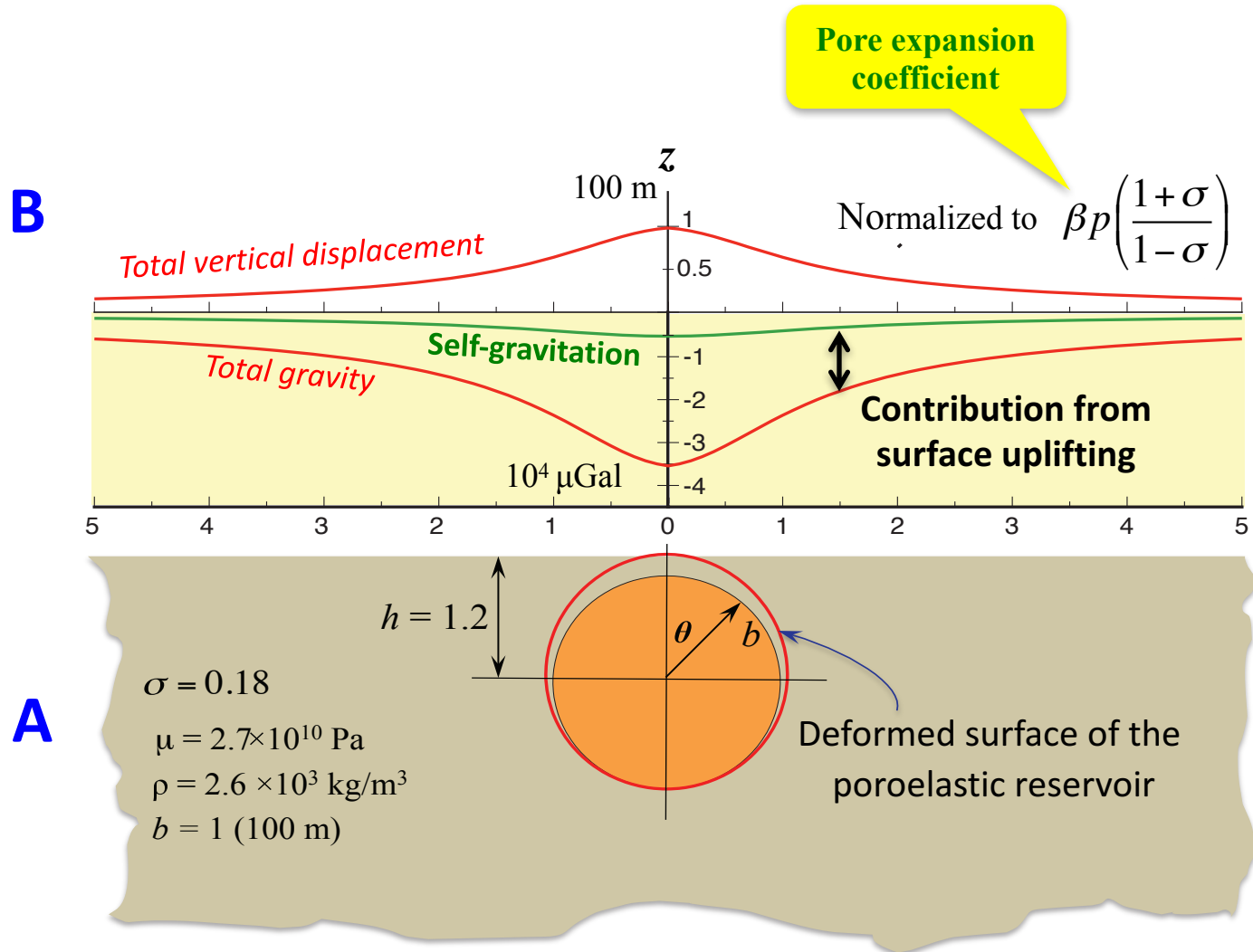
$$\lambda \nabla (\nabla \cdot \mathbf{u}) + \mu \nabla \cdot (\nabla \mathbf{u} + \mathbf{u} \nabla) - (3\lambda + 2\mu) \beta \nabla p = 0$$

Gravity and gravitational gradient from self-gravitation on the free surface

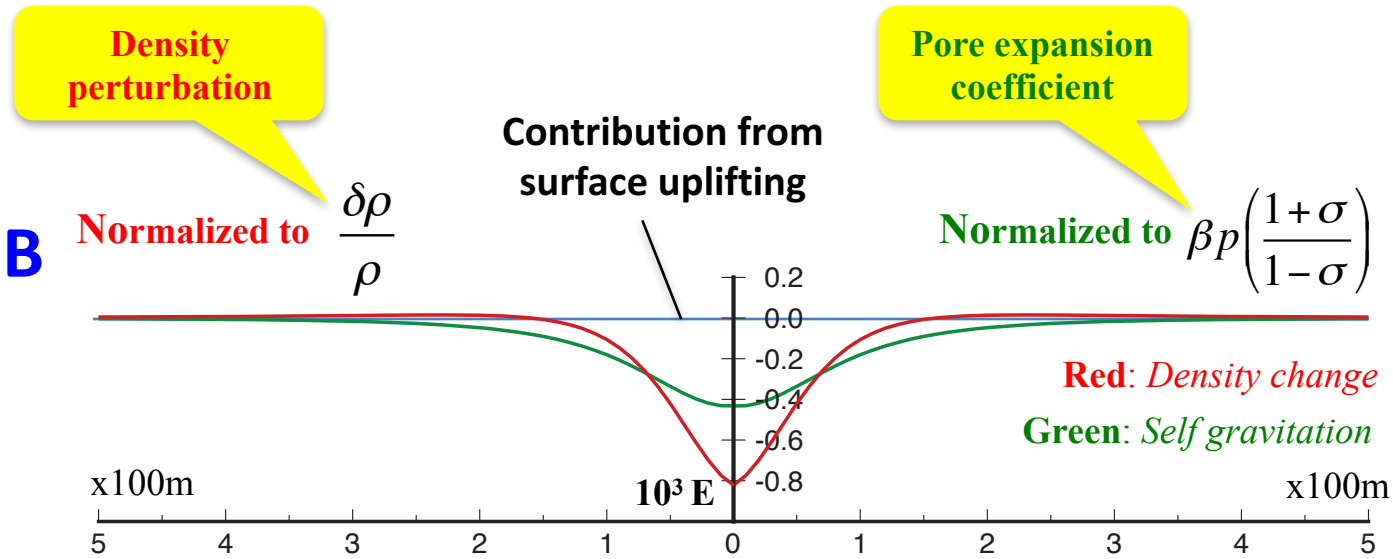
$$-\nabla\phi = -\left(\frac{\partial\phi}{\partial x}, \quad \frac{\partial\phi}{\partial y}, \quad \frac{\partial\phi}{\partial z}\right)$$

$$-\nabla\nabla\phi = -\begin{pmatrix} \frac{\partial^2\phi}{\partial x^2} & \frac{\partial^2\phi}{\partial x\partial y} & \frac{\partial^2\phi}{\partial x\partial z} \\ \frac{\partial^2\phi}{\partial x\partial y} & \frac{\partial^2\phi}{\partial y^2} & \frac{\partial^2\phi}{\partial y\partial z} \\ \frac{\partial^2\phi}{\partial x\partial z} & \frac{\partial^2\phi}{\partial y\partial z} & \frac{\partial^2\phi}{\partial z^2} \end{pmatrix}$$

Perturbed gravity



Perturbed gravity gradient $\frac{\partial g_z}{\partial z}$



A

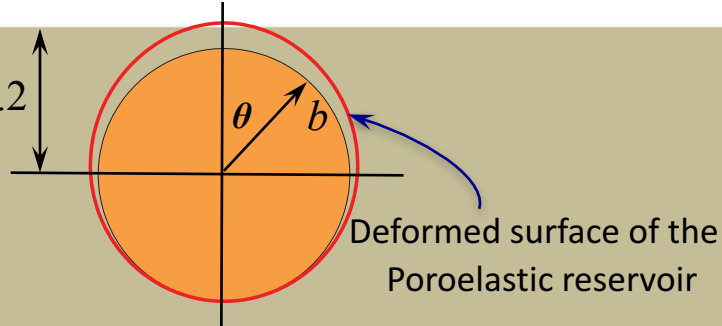
$$\sigma = 0.18$$

$$\mu = 2.7 \times 10^{10} \text{ Pa}$$

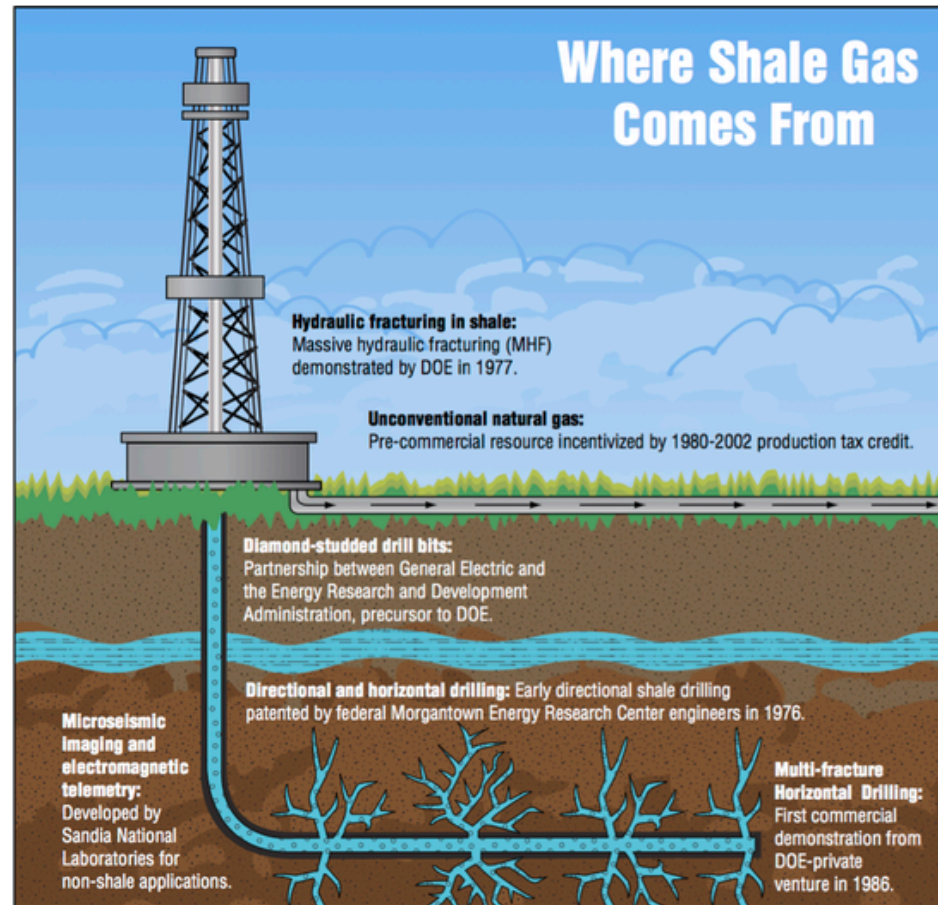
$$\rho = 2.6 \times 10^3 \text{ kg/m}^3$$

$$b = 1 \text{ (100 m)}$$

$$h = 1.2$$

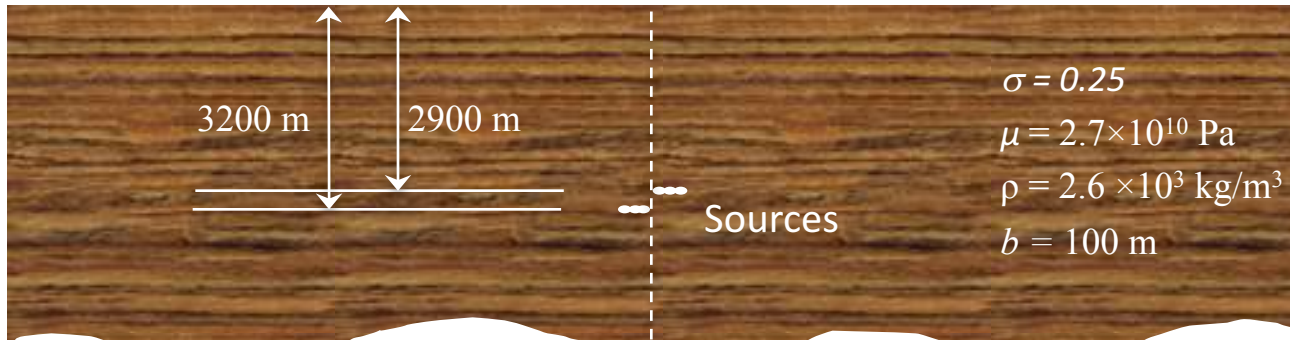
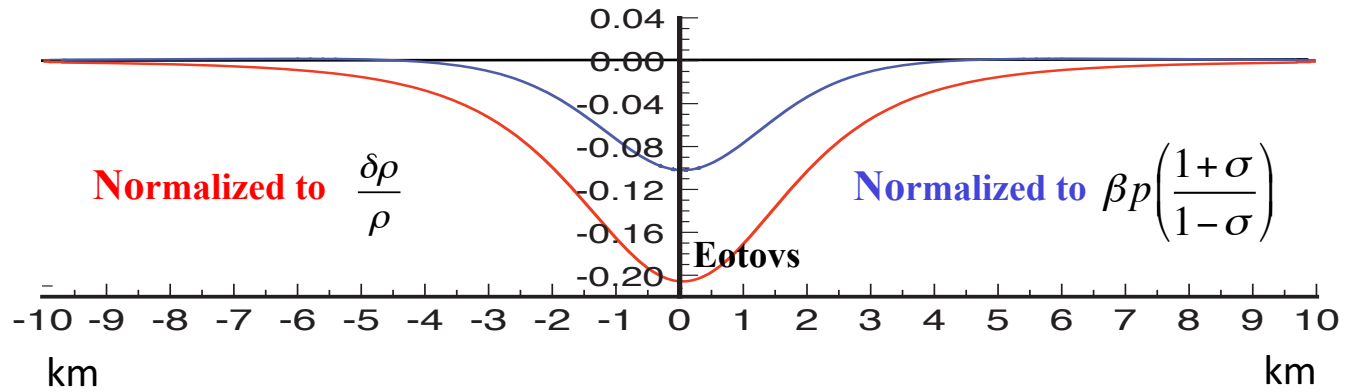


Hydraulic fracturing for shale gas



<https://thebreakthrough.org/index.php/programs/energy-and-climate/where-the-shale-gas-revolution-came-from>

Perturbed gravity gradient from density (red) self-gravitation (blue) and surface uplift (black) $\frac{\partial g_z}{\partial z}$



Conclusions

The gravity gradient can effectively denoise the ground subsidence/uplift effect in deformation-related time-varying gravity variations.